

# Statistical Downscaling of Sea Surface Temperature Projections with a Multivariate Gaussian Process Model

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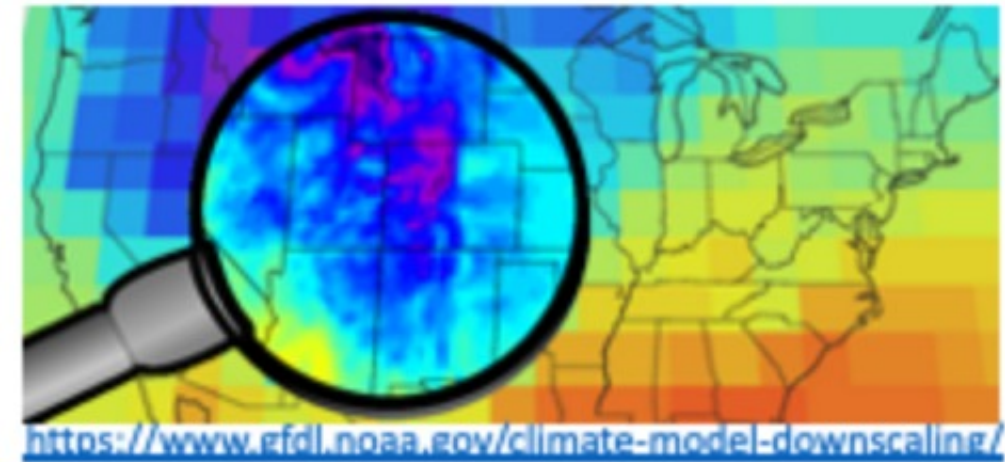
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## Introduction

- Global Climate Models (GCMs) produce projections for future climate.
- GCM projections are usually made at coarse spatial resolutions ( $\sim 1^\circ$ ).



- Downscaled GCM projections are desired for regional studies.
- There are two types of downscaling for climate model outputs.
  - Dynamical downscaling (DD)
    - Use Regional Climate Models (RCMs) and boundary conditions are generated from GCMs.
    - Computationally expensive.
  - Statistical Downscaling (SD)
    - Establish a statistical relationship between coarse-resolution model projections and fine-resolution observations.
    - Computationally less expensive.
- We propose an improved, computationally efficient SD method with a multivariate Gaussian process model.
- We present a demonstration using Sea Surface Temperature (SST).

## Model

- Let  $Y_{t,1}(s)$  be the MUR observational SST at location  $s$  on month  $t$  and  $Y_{t,2}(s)$  be the deterministically downscaled SST from the standard SD method.

- We assume the bi-variate process,  $\mathbf{Y}_t(s) = (Y_{t,1}(s), Y_{t,2}(s))'$  follow,

$$\mathbf{Y}_t(s) = \boldsymbol{\mu}_t(s) + \mathbf{v}_t(s)$$

- $\boldsymbol{\mu}_t(s)$  is the mean component to describe variation with spatio-temporal interaction and  $\mathbf{v}_t(s)$  is a multivariate Gaussian process to capture small-scale spatial variations. We assume,

$$v_{t,i}(s) = \sum_{k=1}^K \xi_i^k U^k(s) + \sum_{l=1}^L \eta_{t,i}^l S^l(s) + \varepsilon_{t,i}(s) \text{ for } i = 1, 2$$

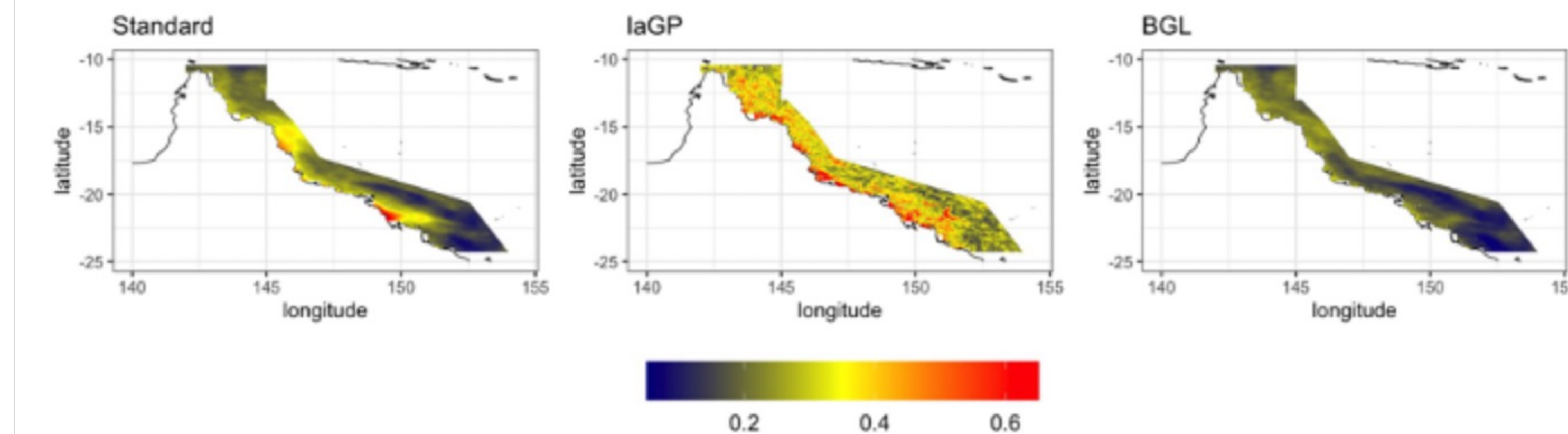
- $\xi_i^k$  for  $k = 1, \dots, K$  are a set of fixed effects and  $\eta_{t,i}^l$  for  $l = 1, \dots, L$  are a set of random effects to describe additional spatial variability but assumed to be independent across time.
- Motivated by Basis Graphical Lasso (BGL) we set  $U^k(s)$  and  $S^l(s)$  to be known orthogonal basis functions (Krock et al., 2021).
- Let  $\boldsymbol{\eta}_t^l = (\eta_{t,1}^l, \eta_{t,2}^l)'$  for  $l = 1, \dots, L$  and assume,  $\boldsymbol{\eta}_t^l \sim N(\mathbf{0}, \mathbf{Q}_l^{-1})$ .
- We assume the precision matrix  $\mathbf{Q}$  of the full vector  $\boldsymbol{\eta}_t = (\boldsymbol{\eta}_t^1, \dots, \boldsymbol{\eta}_t^L)'$  is,  $\mathbf{Q} = \text{diag}(\mathbf{Q}_1, \dots, \mathbf{Q}_L)$ .
- Last term  $\varepsilon_{t,i}(s)$  is white noise with zero mean and  $\tau_i^2$  variance which is independent from random effects.

## Results

- We compare Mean Square Error (MSE) of our proposed method with,

- Interpolated raw GCM projections
- State-of-art (Standard) SD method for SSTs (Van Hooidek et al., 2015)
- local approximate Gaussian Process (laGP) (Gramacy et al., 2015)

ssp126				
Season	GCM	Standard	laGP	BGL
Summer	0.396	0.305	0.419	<b>0.297</b>
Autumn	0.472	0.105	0.221	<b>0.088</b>
Winter	0.878	0.250	0.335	<b>0.163</b>
Spring	0.380	0.212	0.356	<b>0.153</b>
Overall	0.531	0.218	0.353	<b>0.175</b>



## Data

- We choose Great Barrier Reef (GBR) region as the study region.
- We use,
  - Monthly averaged NASA/JPL Multiscale Ultrahigh Resolution (MUR) satellite SST data at 1km resolution from June 2002 to December 2020.
  - Monthly SST outputs from 19 Coupled Model Intercomparison Project Phase 6 (CMIP6) GCM models at 100km resolution under ssp126 climate scenario.
  - ssp 126 is an extreme mitigation scenario which assumes less warming towards next century.
  - Our choice of the scenario here is only for the demonstration purpose.



<http://www.mapnavy.com/map-of/Great-Barrier-Reef-Marine-Park-map>

## Implementation

- Let  $\mathbf{U}(s) = (U^1(s), \dots, U^K(s))'$  and  $\mathbf{S}(s) = (S^1(s), \dots, S^L(s))'$ .

- The model for the full data vector can be written as,

$$\begin{pmatrix} Y_t(s_1) \\ \vdots \\ Y_t(s_N) \end{pmatrix} = \begin{pmatrix} \mu_t(s_1) \\ \vdots \\ \mu_t(s_N) \end{pmatrix} + \begin{pmatrix} \mathbf{U}'(s_1) \otimes \mathbf{I}_2 \\ \vdots \\ \mathbf{U}'(s_N) \otimes \mathbf{I}_2 \end{pmatrix} \begin{pmatrix} \xi^1 \\ \vdots \\ \xi^K \end{pmatrix} + \begin{pmatrix} \mathbf{S}'(s_1) \otimes \mathbf{I}_2 \\ \vdots \\ \mathbf{S}'(s_N) \otimes \mathbf{I}_2 \end{pmatrix} \begin{pmatrix} \boldsymbol{\eta}_t^1 \\ \vdots \\ \boldsymbol{\eta}_t^L \end{pmatrix} + \begin{pmatrix} \varepsilon_t(s_1) \\ \vdots \\ \varepsilon_t(s_N) \end{pmatrix}$$

where,  $\boldsymbol{\xi}^k = (\xi_1^k, \xi_2^k)'$  and  $\boldsymbol{\eta}_t^l = (\eta_{t,1}^l, \eta_{t,2}^l)'$ .

- Estimation is done in two-steps.

- Step1: Detrend and obtain the detail residuals vector  $\mathbf{Z}_t = \mathbf{Y}_t - \hat{\boldsymbol{\mu}}_t - \mathbf{U}\hat{\boldsymbol{\xi}}$ . Then the log-likelihood for the residuals is written as,

$$\log(\det(\boldsymbol{\Sigma})) + \frac{\sum_{t=1}^T \mathbf{Z}_t' \boldsymbol{\Sigma}^{-1} \mathbf{Z}_t}{T} = \log(\det(\boldsymbol{\Sigma})) + \text{tr}(\hat{\boldsymbol{\Sigma}} \boldsymbol{\Sigma}^{-1}) \text{ where, } \text{var}(\mathbf{Z}_t) = \boldsymbol{\Sigma}, \hat{\boldsymbol{\Sigma}} = \frac{\sum_{t=1}^T \mathbf{Z}_t \mathbf{Z}_t'}{T}, \boldsymbol{\Sigma} = \mathbf{S} \mathbf{Q}^{-1} \mathbf{S}' + \mathbf{D} \text{ and } \mathbf{D} = \text{diag}(\tau_1^2, \tau_2^2) \otimes \mathbf{I}_N$$

- Step2: Solve for,

$$\hat{\mathbf{Q}} \in \arg \min \log(\det(\mathbf{S} \mathbf{Q}^{-1} \mathbf{S}' + \mathbf{D})) + \text{tr}(\hat{\boldsymbol{\Sigma}} (\mathbf{S} \mathbf{Q}^{-1} \mathbf{S}' + \mathbf{D})^{-1}) + P(\mathbf{Q}) \text{ where,}$$

$P(\mathbf{Q}) = P(\mathbf{Q}_1, \dots, \mathbf{Q}_L) = \rho \sum_{l=1}^L \sum_{i \neq j} |(\mathbf{Q}_l)_{ij} - (\mathbf{Q}_{l+1})_{ij}|$  and  $\rho$  is a penalty which penalize  $\mathbf{Q}_l$  matrices at adjacent levels if the off-diagonals are not similar (Krock et al., 2021).

## Conclusions and Future Work

- Our downscaling method is computationally feasible for large data sets, accounts for spatio-temporal dependencies, provides meaningful uncertainty estimates, and produce improved downscaling results compared with the state-of-the-art methods.
  - Downscaling for the GBR region (a total of 309,700 high-resolution pixels) was performed in Matlab on a Macbook Air with an 8-core 3.2GHz processor and 8 GB RAM.
  - Given the second process  $Y_{t,2}(s)$  for  $t > T_0$  for all  $s \in \mathcal{D}$ , computation time to fit the model and estimate SSTs for a single future month is 6.8 seconds.
- A possible extension is to generalize the current model to the framework of autoregressive co-kriging for multi-fidelity model output and then consider the observations, regional climate model output, and global climate model output as the high-, medium-, and low-fidelity data, respectively.

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